# PRELIMINARY GEOLOGIC DESCRIPTION

**ARCHITECT** 

**Read** 

 $\ddot{\phantom{0}}$ 

MOBIL PRODUCING<br>G. E. CARGILL NO. 14

December, 1986

Prepared by

Shirley P. Dutton, Michael A. Fracasso, and Robert J. Finley

Bureau of Economic Geology<br>W. L. Fisher, Director<br>The University of Texas at Austin<br>Austin, Texas 78713-7508

Prepared for

The Gas Research Institute<br>Contract No. 5082-211-0708<br>Robert J. Finley - Principal Investigator

 $\ddot{\phantom{a}}$ 

## INTRODUCTION

I

**I** 

I

**INGLES** 

**I** 

I

I

I

**INCHES** 

**INSTEAD** 

**Indian** 

**I** 

en 1999.<br>In de la Galeria

I

I

I

I

I

**Index** 

Two intervals of the Travis Peak Formation were cored in the Mobil Producing Cargill No. 14 well, Waskom field, Harrison County, Texas. Core was recovered from 5,903.0 to 5,960.4 ft and from  $6,148.0$  to  $6,297.1$  ft. The top of the Travis Peak is at 5,901 ft (log depth), so the core begins close to the top of the formation.

# MACROSCOPIC CORE DESCRIPTION

The Travis Peak cores were described using a hand lens and binocular microscope, and graphic logs of the cores were made at a scale of 1 inch = 5 ft I (fig. 1). The following features of the cores were noted on the descriptive I logs: depth; rock type; accessories such as pyrite, organic matter, and burrows; sedimentary structures; texture (sorting); induration; grain size; relative amount of carbonate cement; color; and special features such as reservoir bitumen (solid hydrocarbons) and calcareous nodules (fig. 1). Porosity and permeability values reported by Reservoirs, Ind. are noted on the graphic logs, as are the depths from which thin sections have been made. Correction factors between core and log depths are noted for each cored interval (fig. 1). The cores consist of fine to very fine sand stone, mudstone, and limestone. The two cored intervals are described below.

Sandstones in the lower core are interbedded with thick red and gray I  ${\small \texttt{mudstones.}}$  The sandstones generally have sharp lower contacts and may have I concentrations of carbonate nodules or ripped  $\mu$  mud clasts at the base. The most abundant sedimentary structures are planar to slightly inclined laminae and ripples; high-angle crossbedding is not common. The lack of abundant crossbeds suggests that either the current energy was moderate to low or water

depth was shallow, or both. Burrowing increases in abundance in the sandstones higher in the cored interval. The thick, lowest sandstone  $(6,234$  to  $6,280$  ft, core depth) is burrowed only at the top, whereas the middle sandstone  $(6,188)$  to 6,205 ft) is moderately burrowed throughout, and the upper sandstone  $(6,151)$  to 6,169 ft) is extensively burrowed at the top  $(fig. 1)$ .

I

I

**INSTEAD OF** 

**Inches** 

I

**INSTERNE** 

In the case of

I

I

I

I

**INSTEAD OF** 

**Index** 

I

I

I

[

 $\blacksquare$ 

I

Mudstones in the lower cored interval commonly contain calcareous nodules. Pyrite and wood fragments occur in some of the gray mudstones, particularly at 6,210 ft. Root traces were observed in mottled gray and red mudstone from 6,169 to 6,175 ft (fig. 1).

Rocks of the lower core probably are alluvial or delta-plain deposits; , there is no evidence of marine deposition in this interval. Sandstones represent small fluvial channels or crevasse splays deposited in the adjacent floodplain. Mudstones are relatively thick and probably were deposited in floodplain environments, including well-drained and poorly drained swamps. The gray color of some of the mudstones was probably caused by a high water table, abundant organic matter, and reducing fluids that moved through the overlying sandstones.

Sandstones in the upper cored interval are thinner than those in the lower zone; they generally have sharp bases and fine upward. Ripples, planar to slightly inclined laminae, contorted beds, and crossbeds are the most common I sedimentary structures. Thin sandstones and the tops of thicker sandstones are commonly burrowed. Thinly interbedded sandstone or siltstone with mudstone is common (for example, 5,930 to 5,937 ft) betwee<mark>n the thicker sandstones. Mud-</mark> stones at 5,957 ft to 5,959 ft contain abundant carbonate. Oyster shell fragments occur in zones near the top and bottom of the core (fig. 1).

The upper core interval is interpreted as having been deposited in a ! marginal marine position, such as a tidal The thicker, sharp-based

sandstones are probably tidal channels that cut interbedded sandstone and mudstone of the mixed tidal flat. The presence of oysters provides additional evidence for a marine to brackish water depositional environment.

I

**I** 

I

**I** 

I

**I** 

**I** 

**INSTALL** 

**I** 

**Inches** 

**INGLESS** 

 $\label{eq:3.1} \mathcal{G}^{(1)}_{\text{max}}(t)$ 

**INSTRUCTION** 

I

**I** 

I

**[** 

[

I

#### PETROGRAPHIC DESCRIPTION

i I

I

Detailed study of the core is being conducted with a standard petrographic microscope and with a scanning electron microscope (SEM) that includes an I energy dispersive X-ray system (EDX).

#### Grain Size

Analysis of grain size was accomplished by making grain-size point counts of thin sections. Fifty grains per slide were measured along their long dimension, excluding cement overgrowths in order to determine the size of the detrital grains. Mean diameter of sand- and silt<mark>-sized grains was calculated for</mark> each sample (table 1); detrital and authigenic clays were not included in the calculation of mean grain diameter.

Most sand grains are fine or very fine, between 0.063 and 0.25 mm. Most i silt is coarse silt, between 0.031 and 0.063 mm, and clay particles are smaller than  $0.004$  mm. Texturally, the samples are classified as sandstone, silty sandstone, muddy sandstone, sandy mudstone, and sandy claystone (table 1).

#### Mineral Composition

Twenty-two thin sections have been point counted for a preliminary description of mineral composition (table 2). The sandstones are mineralogically

very mature and are classified as quartzarenite or subarkose. Quartz comprises 91% to 100% of the essential framework constituents (quartz, feldspar, and rock fragments). Plagioclase feldspar is more abundant than orthoclase, and total feldspar volume varies from 0% to 9.0%. Rock fragments, mainly low-rank metamorphic rock fragments, constitute between 0% and 1.5% of the framework grains.

I

I

**ISBN** 

**I** 

**International** 

**Index** 

I

I

I

I

I

I

I

**I** 

I

I

[

 $\blacksquare$ 

**Industry** 

Authigenic cements constitute between  $\rho$ .5% and 26.5% of the sandstone volume in these five samples (table 2). Authigenic quartz, illite, chlorite, kaolinite, dolomite, ankerite, anhydrite, and $\vert$ reservoir bitumen have all been I observed in Mobil Cargill No. 14 thin sections $\vert \cdot \vert$ 

Authigenic illite and chlorite occur as pore-filling cements in secondary pores that formed by the dissolution of feldspars (fig. 2). These cements formed relatively late in the burial history of the Travis Peak. Kaolinite cement was observed in primary pores in the Mobil Cargill No. 14 samples. Petrographic evidence is equivocal, but the kaolinite appears to have precipitated at about the same time or somewhat after the quartz overgrowths.

Quartz cement (fig. 3) is the most abundant authigenic mineral in the Mobil Cargill No. 14 cores. Quartz overgrow hs fill as much as 18.5% of the sandstone volume (table 2), and precipitation of authigenic quartz occluded much of the primary porosity. However, compared with other parts of the Travis Peak, the volume of quartz overgrowths in the Mobil Cargill No. 14 samples is relatively low, and significant primary porosity has been retained. The sandstone at 5,946 to 5,956 ft has porosity as high as 19% and permeability as I high as 332 md (fig. 1). Primary porosity in a sample from  $5,953.0$  ft can easily be seen in SEM (fig. 4); the volume of quartz overgrowths in this sample is only 12%.

Dolomite and ankerite (fig. 5) both oc $\frac{1}{4}$  ur in the Mobil Cargill No. 14 cores. Ankerite cement has a maximum volume of 8.5% in the sample from

I

5,932.2 ft, and it could cause completion problems (formation of an ironhydroxide gel) if it is treated with acid.

I

**INCL** 

I

**I** 

I

**Contact** 

**Inc.** 

I

I

I

I

I

I

I

I

I

I

**.** 

**I** 

Minor amounts of late anhydrite cement were observed in the sample from 5,948.5 ft. The anhydrite was not intersected on the point-count traverse, and therefore it probably has a total volume of less than  $0.5%$ . The anhydrite cement is a late, diagenetic feature and did not precipitate in the depositional environment.

Reservoir bitumen (solid hydrocarbons) occurs in both the upper and lower cores (fig. 1). Where it is abundant, reservoir bitumen fills primary and secondary porosity. Petrographic evidence suggests it entered the sandstones after precipitation of quartz overgrowths and 4nkerite. Preliminary studies of I the reservoir bitumen show it occurs mainly in the upper Travis Peak (Dutton, 1985). Geochemical data indicate it formed by deasphalting of pooled oil after solution of gas into the oil (Rogers and others, 1974).

## Porosity

Porosity observed in thin section varies from  $0.5%$  to  $14.5%$  (table 2). Both primary and secondary pores are present (fig. 5). Secondary pores are formed by the dissolution of framework grains, so they are approximately the same size as detrital grains. However, they c<mark>ommonly contain authigenic clays</mark> and fragments of dissolved framework grains, particularly feldspar. Porosity measured by point counting thin sections is lower than porosity measured by porosimeter on adjacent samples because of the presence of abundant micro-I porosity. Microporosity occurs in detrital and authigenic clays (fig. 2); such porosity generally cannot be seen in thin section, but it can be observed by SEM and is measured by a porosimeter.



Table 1. Grain-size distribution in Mobil Cargill No. 14.

 $\sim 10^{11}$  km s  $^{-1}$ 

**- ... \_----------------**



Table 2. Petrographic analyses of Mobil Cargill No. 14.

 $\frac{1}{2}$ Zircon Tourmaline

------

# **\_ ..... ----------------**





 $\frac{1}{2}$ Zircon Tourmaline

 $\sim 10^{-1}$ 

**Inches** 

I

I

I

I

I

**I**ssued

**I** 

I

I

I

**I** 

I

**I** 

**I** 

I

**I** 

I

I

- Dutton, S. P., 1985, Petrography and diagenes is of the Travis Peak (Hosston) Formation, East Texas: The University of Texas at Austin, Bureau of Economic Geology, topical report prepared for the Gas Research Institute under contract no. 5082-211-0708.
- Rogers, M. A., McAlary, J. D., and Bailey, N. J. L., 1974, Significance of reservoir bitumens to thermal-maturation studies, western Canada Basin: American Association of Petroelum Geologists, v. 58, no. 9, p. 1806-1824.

## FIGURE CAPTION

Figure 1. Descriptive log of core of the Travis Peak Formation from the Mobil Cargill No. 14 well, Harrison County, Texas. Core depths are 5,903.3 to 5,960.4 ft and 6,148.0 to 6,297.1 ft. :

> , i

















 $\frac{1}{2}$ 



in the adjustice make the

# **EXPLANATION OF SYMBOLS**



Porosity

**STRUCTURES** 

## **TEXTURE**



#### **INDURATION**

- WI Well indurated
- I Indurated
- $IF$ Indurated but friable
- IS Indurated but shaly

#### RELATIVE CALCITE CONTENT

- | Slight effervescence
- 3 Moderate effervescence
- 5 Strong effervescence
- 10 Very strong effervescence

# **COLOR**

Abbreviations from Rock-<br>Color Chart, Geological<br>Society of America

Trough crossbedding Planar crossbedding Crossbeds with oversteepened foresets Indistinct cross-stratification Gently inclined lamination  $= GT$ v v  $\mathbf{v}$ an an 84 'QL 3822

Gently inclined lamination separated by<br>low-angle discordances Horizontal lamination Ripple trough lamination

Planar ripple lamination

Climbing-ripple lamination

**Heavily bioturbated sandstone** 

"Massive" sandstone = M

Contorted bedding

Graded bedding

# TRAVIS PEAK CORE

**I** 

**INSURANCE** 

**I** 

**Indian** 

I

i<br>Isabet

I

I

**INC.** 

**INCORPORA** 

**Industry** 

**I** 

**Inches** 

**Index** 

**I** 

**INST** 

I

[

I

Abbreviations from Geological Society of America Rock-Color Chart used on core descriptions.

# HUE N



# HUE YR

 $\overline{1}$ 



# HUE Y



# HUE G



# HUE R



Travis Peak Core Page 2

I

**Inches** 

**I** 

I

**INSIDE** 

**I** 

I

**I** 

**ISBN** 

I

 $\sim$ 

 $\hat{\mathcal{A}}$ 

 $\overline{\phantom{a}}$ 

**I** 

I

I

I

**Inches** 

I

I

 $\blacksquare$ 

I

HUE GY



# FIGURE 2.

SEM photograph of authigenic chlorite plates and quartz overgrowth (Q). Depth is  $6,162.7$  ft. Bar length is 10  $\mu$ m; magnification is 3500X.

# FIGURE 3.

Photomicrograph of subarkose from a depth of 6,245.0 ft. Primary pores have been completely occluded by interlocking quartz cement. Long dimension of photo =  $2.6$  mm; crossed-polarized light.

## FIGURE 4.

SEM photograph of euhedral quartz overgrowths and primary porosity. Depth is 5,953.0 ft. Bar length is  $100 \mu m$ ; magnification is  $150X$ .

## FIG U R E 5.

Photomicrograph of a large mass of ankerite cement (A). Primary (P) and secondary (S) pores are common. Depth is 5,948.5 ft. Long dimension of photo is 2.6 mm; plane-polarized light.

**-------------------**